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# Meteorological Effects on Ground-levels Ozone Metrics in Bangkok Metropolis Region

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Abstract—Multiple linear regression models were constructed to characterize ground-level O<sub>3</sub> metrics in Bangkok Metropolis Region where meteorological parameters are different from other studies in cold cities.  $SAS^{\oplus}$  9.2 software analyzed 2.9-million hourly data during 1997 – 2011 including  $O_3$ ,  $NO_2$  and meteorological variables such as temperature (T), rainfall (RF), relative humidity (RH), pressure (P), solar radiation (SR), wind speed (WS) and wind direction (WD). The results showed O3 was highest in winter because of clearest sky and an atmospheric inversion. O3 had negatively correlated with RH and RF and positively correlated with SR and previous day O<sub>3</sub> (O<sub>3(d-1)</sub>) Natural logarithm transformed  $O_3$  was used for  $O_3$  metrics (daily average, daily maximum and daytime average) for 4 periods (annual, summer, winter and rainy season). Regression results showed that the  $lnO_{3(d-1)}$  was a main positive predictor and RH is the strongest negative predictor following by a positive SR predictor. In winter, major predictors are RH, NO2, WD and lnO<sub>3(d-1)</sub>. In raining season, P and SR played significant positive predictors. In summer, RH is only a main predictor. For validation analysis, the lnO3 daily maximum and daytime average in summer show the highest  $R^2$  values at 0.573 and 0.568 respectively. This work investigated the effects of Bangkok tropical climate parameters influencing O<sub>3</sub> metrics.

 ${\it Keywords}$ —ozone, meteorology, multiple linear regression, seasonal effects

#### I. Introduction

Ground-level ozone  $(O_3)$  is a secondary pollutant, which is not emitted directly, but it can be formed by complex photochemical reactions in the troposphere. The Thai Pollution Control Department (PCD) has been reporting that hourly  $O_3$  levels in Bangkok and its vicinity have been exceeding both 8-hour and 1-hour standards because of increasing automobile vehicles and urban heat island effect. Traffic pollutants such as hydrocarbons (HC) and oxides of nitrogen  $(NO_x)$  can form  $O_3$  in the presence of sunlight. The tropospheric ozone can negatively affect human health and environment. It reduces visibility when reacting with particulate matters in the atmosphere and forms photochemical

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\*Corresponding author Sitthichok Puangthongthub, Ph.D Department of Environmental Science, Faculty of Science, Chulalongkorn University Bangkok, Thailand smog resulting in adverse respiratory and cardiovascular health effects.

Climate and seasonal changes in meteorological factors have showed links with  $O_3$  fluctuations [1, 2]. The favorable meteorological conditions can lift up O<sub>3</sub> concentrations. Solar radiation is the most important factor in  $O_3$  synthesis [3, 4]. Temperature, a surrogate of solar radiation, and the Peroxy Acetyl Nitrate (PAN), naturally released and acting as a source of NO<sub>2</sub> are also associated with increased O<sub>3</sub> [4, 5. Several studies reveal that temperature and heat island effect are well associated with increased O<sub>3</sub> especially in cities where highrise buildings and properties of constructed surfaces help sink  $O_3$  precursors [6, 7]. Wind speed and direction can dilute  $O_3$ level or concentrate it by transporting it from neighboring cities. In dense urban setting area, wind may not be able to clear the atmospheric completely from air pollutants due to structural characteristic of buildings [8, 9]. Thus the previous day's pollutant concentration is useful in predicting next day's concentration as well as pressure, relative humidity and rainfall are [10, 11]. Several works have applied these metrological variables and O<sub>3</sub> precursors in modeling urban O<sub>3</sub> concentration by using correlation coefficient and multiple linear regression (MLR) analysis [4, 8, 9, 10, 11, 12, 13, 14]. This work aims to investigate the influence of meteorological factors on O<sub>3</sub> concentrations by MLR method in Bangkok where its meteorological condition depends on year-round strong solar radiation and high relative humidity with a presence of monsoon differing from other study locations in cold countries.

## п. Materials and Methods

#### A. Area and Data

This work acquired 2.9 million hourly measurements of  $O_3$ ,  $NO_2$  and meteorological parameters such as temperature (T in °C), solar radiation (SR in MJ/m²), wind speed (WS in m/s), wind direction (WD in degree), relative humidity (RH in %), rainfall (RF in mm) and pressure (P in mmHg) during 1997-2011 from 23 PCD air quality stations in Bangkok metropolis region including surrounding 4 provinces (Pathumthani, Samut Prakarn, Samut Sakhon and Nonthaburi). Ambient air quality monitoring network is showed in Fig. 1. Hourly ozone data were calculated for 3  $O_3$  metrics (daily maximum, daily average, and daytime average of 09.00-17.00 hr.), hourly  $NO_2$ , WS, WD and RH were estimated for daily average, hourly T and the previous day  $O_3$  ( $O_{3(d-1)}$ ) were estimated for daily maximum, and hourly SR and RF were aggregated for daily total.



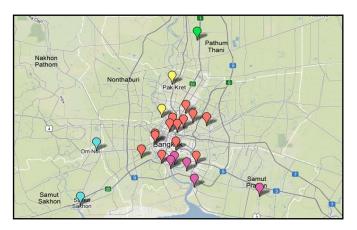


Figure 1. Ambient air quality monitoring stations of PCD in Bangkok Metropolis Regions

# B. Methodology

Pearson product-moment correlation coefficients (r) were computed for 4 weather periods to witness how well each O<sub>3</sub> metric was correlated with its predictors (NO2, T, SR, WS, WD, RH, RF, P and O<sub>3(d-1)</sub>). We then fit 12 MLR models (3 O<sub>3</sub> metrics for 4 sub analyses) to characterize what meteorological factors were annually and seasonally influencing O<sub>3</sub> metrics significantly. The mathematical expression of MLR equation can be written in the form shown in (1).

$$y = a + b_1 x_1 + \dots + b_k x_k \tag{1}$$

where y is participant's predicted scores on the criterion variable (the dependent variable), x\_k is the kth predictor variables (the kth independent variables), a is an intercept constant (the regression constant) and b\_k is the nonstandardized multiple regression coefficient for the kth predictor variables (the kth regression coefficient). This study used the stepwise method that is the combination method of backward and forward method to optimize prediction models [15]. Each O<sub>3</sub> metric (y variable) regressed on its predictors (x variables) such as NO<sub>2</sub>, O<sub>3(d-1)</sub> and the meteorological parameters using SAS® 9.2 software.

# ш. Results and Discussions

# A. Temporal exploratory analysis

Seasonal O<sub>3</sub> daily average fluctuations were observed as shown in Fig.2 with a 14-year average at  $15.36 \pm 11.01$  ppb (N = 1,849,697) ranging from few ppb to 56 ppb. The  $O_3$  peaks were in winter at an average of 18.96 ± 20.68 ppb (N= 615,606) following by summer with an average of 17.75  $\pm$ 17.6 ppb (N = 44,3630) and rainy with an average of 10.97  $\pm 17.16$  ppb (N = 788,121). Winter O<sub>3</sub> levels were highest but less fluctuating than summer O<sub>3</sub> levels because of less cloud with strong radiation and shorter atmospheric mixing height for well promoting photochemical reaction of O<sub>3</sub> precursors while their temperature levels were not much different i.e.,  $27.92 \pm 3.27$  °C vs.  $30.01 \pm 3.00$  °C respectively. The lowest O<sub>3</sub> average found in rainy season was likely due to more cloudy days resulting in low solar radiation and wet deposition (RF and RH) of O<sub>3</sub> precursors [18].

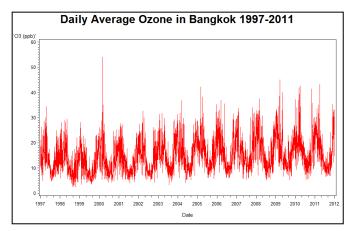


Figure 2. Daily average ozone concentrations from 23 PCD air quality stations in Bangkok metropolis region during 1997 to 2011

# **B.** Correlation Coefficient Analysis

Most correlation coefficients were found statistically significant (P<0.05) except few indicated with star symbol as shown in Table 1. NO<sub>2</sub> levels were positively correlated with O<sub>3</sub> maximum in all tests but negatively correlated with other two metrics in 3 seasons likely due to natural characteristic of unstable species of NOx and O3 precursor mixing speed under different meteorological conditions. The  $O_{3(d-1)}$  concentrations were most strongly positive (r ranging from 0.57 to 0.69) in all periods due to day-to-day accumulation [10, 11]. In all periods, positive correlations were observed for SR and negative correlations were seen for RH and RF consistently. For T, O<sub>3</sub> maximum and daytime average (two O<sub>3</sub> metrics during solar radiation available) showed consistent positive correlation but for WS, they had negative correlation consistently. Pressure trended to be positively correlated in many tests, i.e. high P promoted well O<sub>3</sub> precursor mixing except few tests in summer with negligible r values. Among meteorological parameters, RH was predominantly and negatively correlated (r average at -0.27) and associated with rainy days when cloudier sky and lower SR minimize photochemical production while wet deposition diluting O<sub>3</sub> precursors happened [4, 8] following by SR positively correlated (r average at 0.18).

Summer O<sub>3</sub> metrics showed strong positive correlation with SR and T but strong negative correlation with RH. Previous studies demonstrated O<sub>3</sub> concentrations were high under high T, strong SR and low RH [4, 9, 19]. In rainy season, T, SR and P were in positive correlation with all O<sub>3</sub> metrics and in opposite direction for RF, RH and WD. In winter, we found SR, WD and P showed positive correlation but RF and RH showed negative correlation. Although in rainy season RH was high and expected to have high negative correlation coefficient but we saw this correlation in summer and winter instead. This may be due to high fluctuation of RH between wet and dry days comparing resulting in large SD of O<sub>3</sub> daily average. For solar radiation, it was positive in all

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tests as tropospheric  $O_3$  are well produced during appearance of strong solar radiation.

## c. Multiple Linear Regression Analysis

The natural logarithm transformation used for all O<sub>3</sub> metrics has improved model R<sup>2</sup> (R<sup>2</sup> results of non-transformed O<sub>3</sub> were not shown). Multicollinearity (by variance inflation factor, VIF) and tolerance statistics were also analyzed showing no multicollinearity among predictors (result not shown). The lnO<sub>3</sub> daytime average models showed highest R<sup>2</sup> values in all periods possibly that we modeled O<sub>3</sub> data set only during photochemical period (9-17 hr), following by the lnO<sub>3</sub> daily average and lnO<sub>3</sub>daily max models (see Table 2). The model R<sup>2</sup> values ranged from 0.5019-0.6207 for lnO<sub>3</sub> daytime average, 0.4823-0.5888 for lnO<sub>3</sub> daily average and 0.4823 -0.5677 for  $lnO_3$  daily maximum. The  $lnO_{3(d-1)}$  was robust in all models as a main predictor (regression coefficients (βs) ranging from 0.608- 0.696 ) which is consistent with the similar analysis done in Greater Athens, Greece [10]. NO<sub>2</sub> was a negative predictor for lnO3 daily and daytime average metrics in all periods. This relationship was expected because NO<sub>2</sub> was an O<sub>3</sub> precursor and was decreased to from O<sub>3</sub> [20]. However this was not seen in most lnO<sub>3</sub> daily maximum models that predicted only an hour with the highest O<sub>3</sub> so 24hour average of NO<sub>2</sub> may not be an effective predictor for this case.

For the meteorological parameters, RH is the strongest negative predictor following by a positive SR predictor. Bangkok has tropical climate with long range of monsoon (6 months). High RH and wet deposition can absorb O<sub>3</sub> that is soluble [8, 18, 21] so rainfall can make O<sub>3</sub> levels

lower in the atmosphere [6, 20]. Long period of SR can result in adding  $O_3$  peak due to the photochemical process [14]. WS appeared to negatively predict  $lnO_3$  daily maximum and daytime average or WS help dilute  $O_3$  in daytime during the presence of SR by wind transportation [22, 23] but during the longer pe

riod covering day and night time, WS can promote mixing of O<sub>3</sub> precursors or help transport O<sub>3</sub> from other vicinity area [14] such as from Samut Prakkarn where the PCD has been reported that O<sub>3</sub> keeps violating the 1-hr and 8hr standards due to additional O<sub>3</sub> precursors from industrial sources. T (max) was seen as a positive predictor only in lnO<sub>3</sub> daily maximum models in all periods as high T causes convection to enhance vertical O<sub>3</sub> transport and causes the photolysis of PAN chemistry leading to more NO2 formed 4, 9]. However in this work, T (max) showed random effects in other two lnO<sub>3</sub> metrics with extended hours of O<sub>3</sub> in averaging or T (max) may not be a well predictor in Bangkok as temperature levels were not much fluctuating year-round unlike many studies in cold cities showing large temperature gradient between seasons where T can be a significant predictor [22, 23].

For season specific effect, we observed consistent high regression coefficients ( $\beta$ s) in winter for RH and NO<sub>2</sub> as negative predictors and WD and lnO<sub>3(d-1)</sub> as positive predictors in all lnO<sub>3</sub> metrics while SR was positively high in both winter and rainy seasons. Winter meteorological parameters of Bangkok are favorable for O<sub>3</sub> formation as lowest RH for less wet deposition of O<sub>3</sub> precursors and O<sub>3</sub>, highest and ready NO<sub>2</sub> to switch to O<sub>3</sub> due to atmospheric inversion, clearest sky for no SR interruption with more extended hours than those

TABLE I. PEARSON PRODUCT-MOMENT CORRELATION COEFFICIENTS BETWEEN O<sub>3</sub> METRICS AND THEIR PREDICTORS BY SEASONAL AND ANNUAL DATA SET

O <sub>3</sub> metrics	O <sub>3</sub>	$NO_2$	P	Rain (total)	RH	T (max)	WD	ws	SR (total)	O <sub>3 max (d-1)</sub>
(a) Summer					•		•			
Daily avg	1	-0.0931	-0.0009	-0.0768	-0.2664	0.0655*	0.0149	0.1444	0.1706	0.5684
Daily max	1	0.1557	0.0028	-0.0349	-0.2286	0.1561	0.0072	-0.0902	0.0735	0.5806
Daytime avg	1	-0.0867	-0.0002	-0.0644	-0.3510	0.1708	-0.0152	-0.0176	0.1767	0.5992
(b) Rainy										
Daily avg	1	-0.1477	0.0534*	-0.0378	-0.1704	0.1822	-0.0144	0.0973	0.2295	0.5680
Daily max	1	0.0389	0.0586*	-0.0206	-0.0635	0.2020*	-0.0406	-0.0650	0.1328	0.5954
Daytime avg	1	-0.1652	0.0462*	-0.0642	-0.2415	0.2411	-0.0036	-0.0199	0.2503	0.6029
(c) Winter										
Daily avg	1	-0.1520	0.0238	-0.1158	-0.3160	-0.0663	0.0584	0.0924	0.2747	0.6225
Daily max	1	0.0353	0.0193	-0.1010	-0.2533	0.0326	0.1655	-0.0674	0.1890	0.6591
Daytime avg	1	-0.2310	0.0228	-0.0997	-0.3830	0.0129	0.0776*	-0.0054	0.2898	0.6767
(d) Annual										
Daily avg	1	0.0325	0.0235	-0.1051	-0.3395	-0.0016	-0.0908	0.0967	0.1629	0.6514
Daily max	1	0.2046	0.0248	-0.0834	-0.2773	0.0433	-0.0434	-0.0766	0.0671	0.6731
Daytime avg	1	-0.0586	0.0319	-0.0936	-0.4093	0.0409	-0.1238*	-0.0373	0.1624	0.6916

<sup>\*</sup> Few coefficients not statistically significant at  $\alpha = 0.05\,$ 



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studies in cold climate countries and different WD possibly promoting  $O_3$  precursor mixing. In raining season, we found regression coefficients of P and SR showed high values whose gradients may be large between wet and dry days thus can clearly be detected by regression as major positive predictors in raining season. In summer, we did not see any predictors showing significant effects except RH. RH was shifting mostly in winter following by summer and raining season respectively. So this RH fluctuating can be a significant predictor and observed through its regression coefficient

# D. Validation of the Models

We estimated the coefficient of determination  $R^2$  values in all 12 models to see how well observed  $O_3$  and predicted  $O_3$  were fit using 2009 data set. The  $R^2$  ranged from 0.3057 to 0.5732 (averaged at 0.4628). In rainy, winter and annual tests, all  $InO_3$  daily average and daytime average had higher  $R^2$  values consistently than those of  $InO_3$  daily maximum. However, in summer the  $InO_3$  daily maximum model showed the highest  $R^2$  of 0.5732 following by the  $InO_3$  daily average model with  $R^2$  of 0.5676 (as seen in Figs 3 and 4 respectively). We also calculated  $R^2$  values for non In-transformed models and their results revealed that the  $R^2$  values of In-transformed In3 models were overall higher than the In4 values of In5 non-transformed In5 models (the highest In8 in daily average metrics in summer at 0.4922 and rainy season at 0.4125).

## **IV.** Conclusion

We analyzed 2.9 million hourly measurements of  $O_3$ ,  $NO_2$  and meteorological parameters in Bangkok and nearby 4 provinces and found positive correlation for SR and  $O_{3(d-1)}$ . The negative correlation was seen for RH and RF. For T, two

 $O_3$  metrics during sunlight showed positive correlation but for WS, they had negative correlation. RH was predominantly and negatively correlated following by SR that was positively correlated. The natural logarithm transformation of  $O_3$  metrics improved model  $R^2$ . The  $lnO_3$  daytime average models showed highest  $R^2$  values in all periods. The  $lnO_{3(d-1)}$  was a major predictor.  $NO_2$  was a negative predictor for  $lnO_3$  daily and daytime average metrics. RH is the strongest negative predictor following by a positive SR predictor. Bangkok has tropical weather with extended hours of SR. WS appeared to be a negatively predictor, not only helping  $O_3$  dilution in daytime but also can promote mixing of O3 precursors. T (max) may not be a well predictor in Bangkok as temperature was not much variable differing from cold countries indicating T was their major positive predictor.

In addition, unique results were observed in winter, favorable to  $O_3$  formation, for example lowest RH for less wet deposition, highest and ready  $NO_2$ , clearest sky for no SR interruption with more extended daytime hours than those studies in cold climate countries and different WD promoting  $O_3$  precursor mixing. In raining season, we found P and SR showed high  $\beta$  values and in summer, only RH was only a significant predictor. This work tested the effects of Bangkok tropical climate parameters influencing different  $O_3$  metrics in different weather periods.

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TABLE II. STANDARDIZED REGRESSION COEFFICIENTS FROM MULTIPLE LINEAR REGRESSIONS WITH LNO<sub>3</sub> CONCENTRATIONS

InO <sub>3</sub> metrics	Standardized regression coefficients										
	$NO_2$	P	Rain	RH	T	WD	WS	SR	$lnO_{3(d-1)}$	$\mathbb{R}^2$	
(a) Summer											
Daily avg	-0.1214	0.0131	-0.0254	-0.1918	-0.0405	0.0238	0.0807	0.0603	0.6082	0.4823	
Daily max	0.0650	0.0193	-	-0.1014	0.0741	0.0133	-0.0383	0.0279	0.6325	0.4673	
Daytime avg	-0.0814	0.0150	-0.016	-0.2134	0.0132	-	-0.0174	0.0574	0.6075	0.5019	
(b) Rainy Season											
Daily avg	-0.1205	0.0666	-	-0.1341	-	-0.0288	0.0406	0.1096	0.6222	0.4989	
Daily max	0.0092	0.0408	-	-0.0303	0.1029	-0.0082	-0.0449	0.0583	0.6555	0.4836	
Daytime avg	-0.1078	0.0417	-0.0206	-0.1366	0.0345	-0.0119	-0.0436	0.1088	0.6280	0.5294	
(c) Winter											
Daily avg	-0.1734	0.0130	-0.0516	-0.2202	-0.0226	0.0548	0.02587	0.0957	0.6256	0.5888	
Daily max	-0.0175	0.0185	-0.0345	-0.1388	0.0285	0.0963	-0.0550	0.0789	0.6683	0.5671	
Daytime avg	-0.1478	0.0138	-0.0278	-0.2152	-	0.0739	-0.0478	0.0831	0.6288	0.6207	
(d) Annual											
Daily avg	-0.0956	0.0178	-0.0229	-0.1867	-0.0335	-0.0199	0.0477	0.0988	0.6888	0.5674	
Daily max	0.0525	0.0172	-0.0136	-0.0985	0.0321	0.0151	-0.0497	0.0650	0.6959	0.5549	
Daytime avg	-0.0852	0.0173	-0.0219	-0.2048	-0.0158	-0.0096	-0.0388	0.0850	0.6635	0.5993	

Using stepwise regression method with significant level at  $\alpha = 0.05$ 



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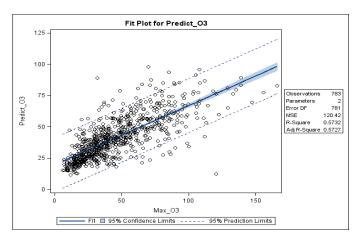


Figure 3. Validation for summer daily maximum O<sub>3</sub> metric using 2009 data

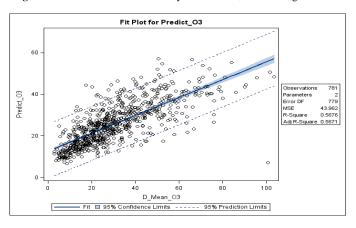


Figure 4. Validation for summer daytime average O<sub>3</sub> metric using 2009 data

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